

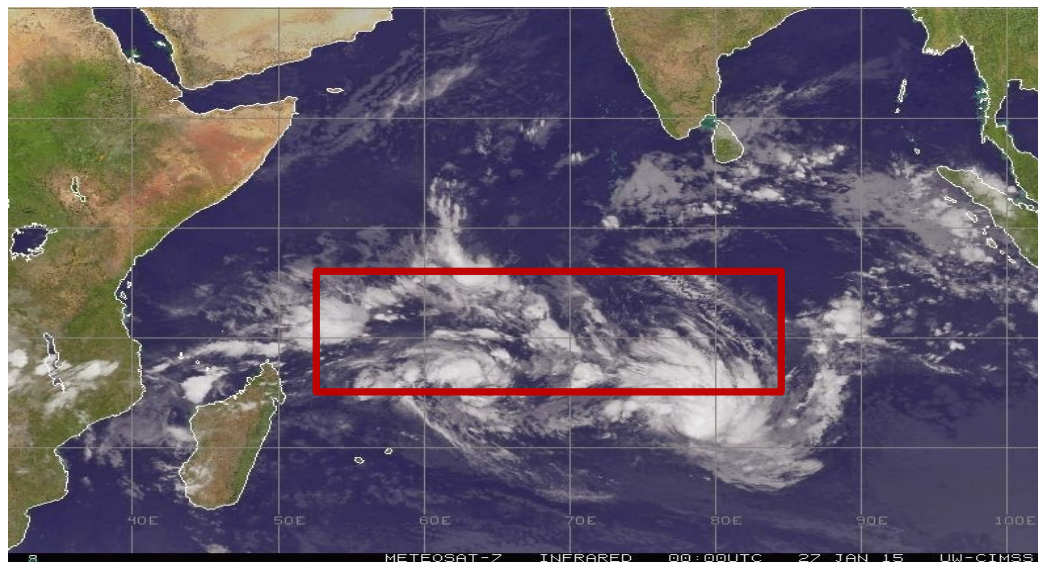
# Upper ocean response to Very Intense Tropical Cyclones in the South West Indian Ocean

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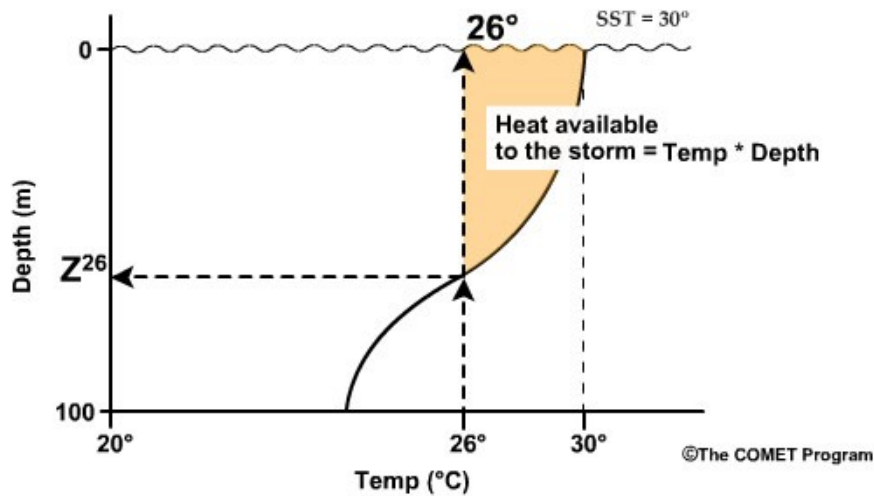


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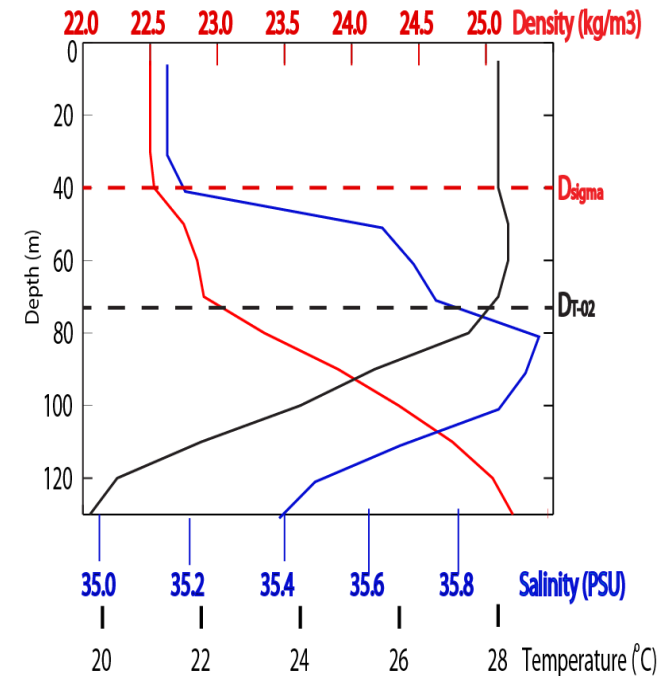
# Brief intro ..

## Tropical Cyclone Heat Potential (TCHP)



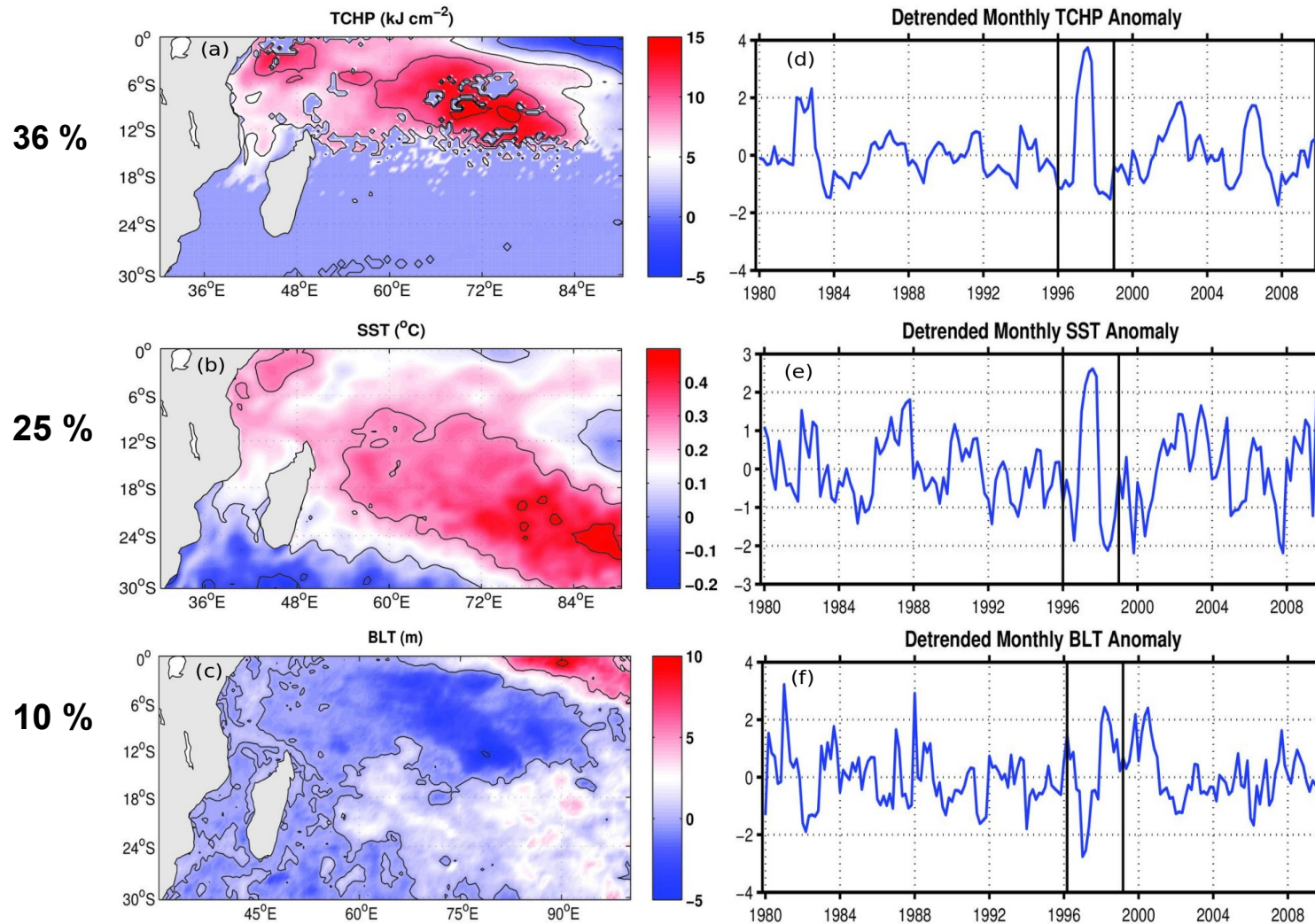
$$TCHP = C_p \int_{z(D_{26})}^{z(0)} (T - 26) \rho dz$$

## Barrier Layer Thickness (BLT)



Halocline shallower than the thermocline, a layer is established between the **MLD** and the **ILD** called the **BLT**

# EOF Analysis



→ TCHP exhibits largest variation in the central and eastern SCTR region ( $r=0.57$  with ENSO)

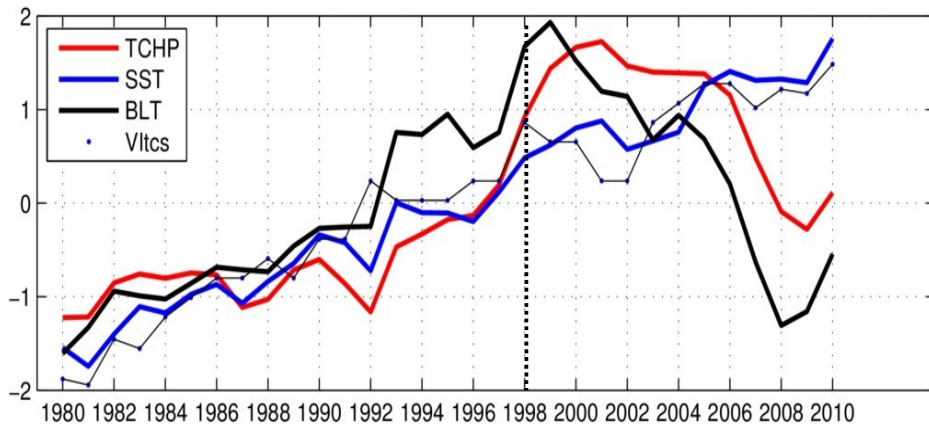
→ SST pattern – typical response to ENSO

→ BLT shows weak loadings over most of the Southern part except in the SE ( $r= 0.53$  with ENSO)

→ SODA reanalysis (monthly output,  $\frac{1}{2}$  deg resolution )

→ Time period : 1980 -2010 – averaged over cyclonic season (Nov - April)

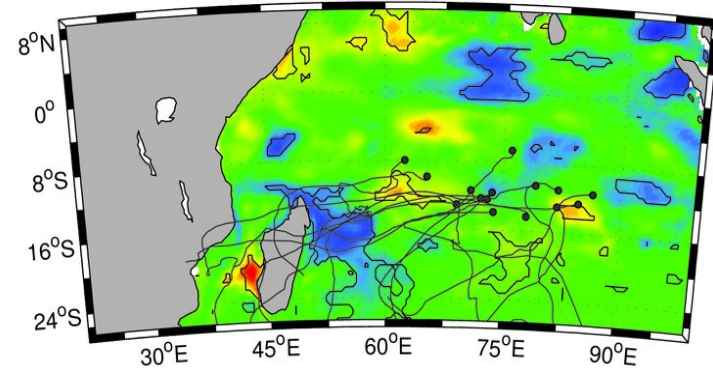
# Variability of TCHP, BLT and Very intense Tropical Cyclone (VITC) intensity



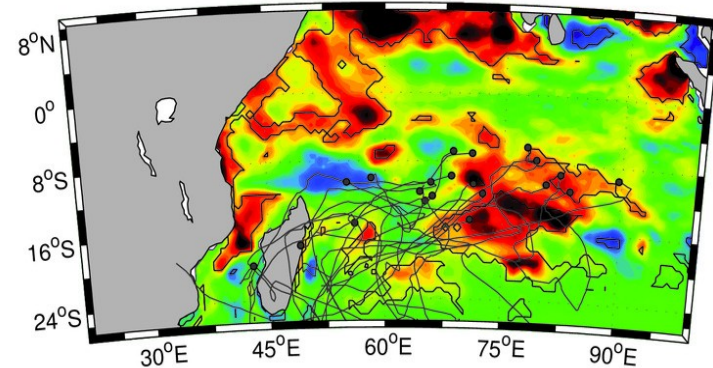
*10-year running through the data, SST, TCHP, BLT and VITCs (Nov-April) across SCTR*

- 1980 -1998, VITCs seem to increase with increase in SST, TCHP and BLT.
- Post 2000, although a decreasing trend in TCHP and BLT is observed, VITCs seem to increase along with a warming in SST.

TCHP trend overlaid with VITC tracks and genesis location 1980–1996

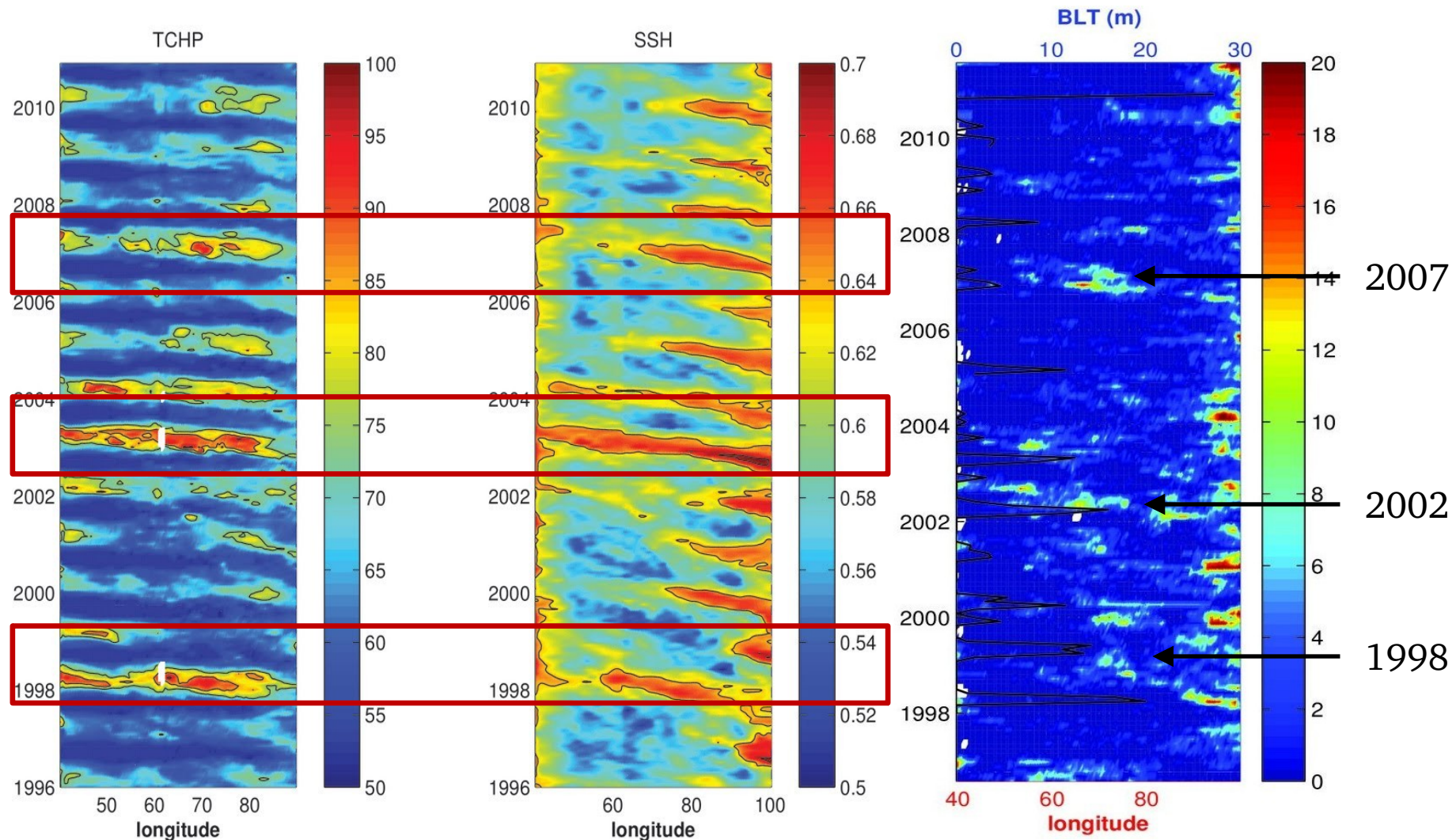


TCHP trend overlaid with VITC tracks and genesis location 1998–2010



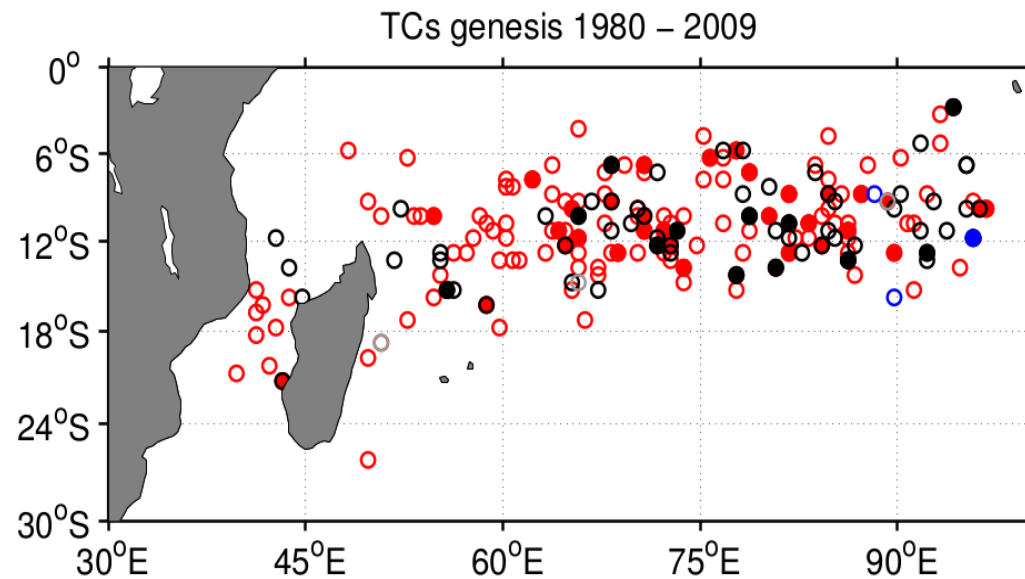
Climate shift after 1998 El Nino event  
– warming over Indian Ocean

# Rossby waves influences



- Westward propagation of rossby waves
- Co-propagation of BLT and TCHP - mechanism of BLT : advection of fresh water from Indonesion throughflow via SEC or heavy precipitation (ITCZ)

# Tdy and BLT influences on TC frequency



*Red : High TDY*

*Black : High TDY, High BLT*

*Blue : High BLT*

*Empty : TCs in general*

*Filled : VITCs only*

|                |                                 | Strong (-ve) | Moderate (-ve)             | Moderate (+ve)                | Strong (+ve)              |
|----------------|---------------------------------|--------------|----------------------------|-------------------------------|---------------------------|
|                | <b>TDY Ano</b><br><b>BL Ano</b> | 0            | 6 - (~ 3%)<br>2 - (~ 4.3%) | 196 - (~88%)<br>41 - (~87.2%) | 20 - (~9%)<br>4 - (~8.5%) |
| Strong (-ve)   | 0                               |              |                            |                               |                           |
| Moderate (-ve) | 153 - (~69%)<br>33 - (~70%)     |              | 3 - (~1,4%)<br>1 - (~2%)   | 137 - (~62%)<br>30 - (~64%)   | 13 - (6%)<br>2 - (2%)     |
| Moderate (+ve) | 44 - (~20%)<br>6 - (~13%)       |              | 3 - (~1,4%)<br>1 - (~2%)   | 39 - (~17%)<br>5 - (~10%)     | 2 - (~0.9%)<br>0          |
| Strong (+ve)   | 25 - (~11%)<br>8 - (~17%)       |              | 0<br>0                     | 20 - (9%)<br>6 - (13%)        | 5 - (~2.3%)<br>2 - (4%)   |

# Conclusion

- Cyclone intensification is largely influenced by large TCHP and deep barrier layer.
- TCHP and BLT are well correlated with El Nino and are modulated by rossby waves.
- Overall warming trend in the Indian Ocean after 1997-1998 El Nino.
- Tropical cyclone genesis is more influenced by temperature than salinity effect.